

INTRODUCTION

The High Plains aquifer underlies parts of Colorado, Kansas, Nebraska, New Mexico, Oklahoma, South Dakota, Texas, and Wyoming. About 20 percent of the irrigated land in the United States is in the region underlain by the High Plains aquifer, and nearly 30 percent of the ground water used for irrigation in the United States is pumped from the High Plains aquifer (Weeks and others, 1988). The large volume of water withdrawn from the aquifer for irrigation purposes since 1940 has had a substantial effect on water levels in the aquifer.

Water-level changes in the High Plains aquifer, however, have not been uniform. Large regional differences in rates of ground-water recharge and withdrawals for irrigation as a result of regional variability in climate, soil, land use, and historical development of irrigated agriculture have substantially affected the geographical patterns of water-level change in the High Plains aquifer.

The High Plains Regional Aquifer-System Analysis, completed by the U.S. Geological Survey (USGS) in the mid-1980's, indicated that substantial water-level declines had occurred in large parts of the High Plains aquifer (Gutentag and others, 1984). Congress reacted that accurate information on the conditions and water-level changes in the High Plains aquifer is necessary to make sound management decisions concerning the use of water, to project future economic conditions, and to conduct hydrologic research pertaining to the High Plains aquifer. Congress passed the Water Resources Research Act of 1984 (Public Law 98-242), which mandated and funded a program for the USGS to monitor water-level changes in the aquifer annually, starting in 1988. The Federal Reports Elimination and Sunset Act of 1995 (Public Law 104-66) and the Omnibus Water Resources Development Act of 1986 (Public Law 98-662) amended Public Law 98-242. Congress now mandates that the USGS in cooperation "...with the States of the High Plains region is authorized and directed to monitor the water levels of the Ogallala (High Plains) aquifer, and report biennially to Congress."

The purpose of this report is to present (1) the water-level changes in the High Plains aquifer for three time periods: predevelopment to nonirrigation season (generally October through March) 1979-80, nonirrigation season, 1979-80 to nonirrigation season 1994-95, and nonirrigation season 1993-94 to nonirrigation season 1994-95; and (2) the precipitation pattern in the region underlain by the High Plains aquifer for 1994 and 1981-94. The water-level changes and precipitation patterns are shown in maps; periodic water levels for selected wells are shown in hydrographs. In this report, the "High Plains region," "central High Plains region," "northern High Plains region," and "southern High Plains region" are the terms used to designate the areas underlain by the High Plains aquifer or its subdivisions (fig. 1).

EXTENT AND DESCRIPTION OF THE HIGH PLAINS AQUIFER

The High Plains aquifer underlies about 174,050 square miles in Colorado, Kansas, Nebraska, New Mexico, Oklahoma, South Dakota, Texas, and Wyoming (fig. 1). The High Plains aquifer, formerly known as the Ogallala aquifer, consists of one or more geologic units of late Tertiary or Quaternary age that are hydraulically connected, including Quaternary deposits, the Ogallala Formation, the Arikaree Group, and the Brule Formation (table 1). The Ogallala Formation is generally the principal unit of the High Plains aquifer (Gutentag and others, 1984).

Table 1. Geologic units of the High Plains aquifer
[Sources: Gutentag and others, 1984; Emery and others, 1987; Tedford and others, 1987; Swinehart, 1989]

Geologic unit comprising the aquifer	System (series) and time of deposition, in millions of years before present	Location where geologic unit is a substantial part of the High Plains aquifer	Colorado	Kansas	Nebraska	New Mexico	Oklahoma	South Dakota	Texas	Wyoming
Dune sand	Quaternary (Holocene), 0.008 to 0.0015	Sand, very fine to medium-grained, unconsolidated	X	X						
Valley-fill and alluvial deposits	Quaternary (Holocene and Pleistocene), 1.5 to present	Clay, silt, sand, and gravel, unconsolidated	X	X		X				
Ogallala Formation	Tertiary (Miocene), 19 to 5	Clay, silt, sand, and gravel, generally cemented by calcareous concretions, mortar beds formed	X	X	X	X	X	X	X	X
Arikaree Group	Tertiary (Oligocene and Miocene), 29 to 19	Sandstone, very fine to fine-grained, with beds of sand, and sandy clay			X			X		X
Brule Formation	Tertiary (Oligocene), 31 to 29	Siltstone, massive, with thin beds of sandstone, calcareous silt, and clay	X		X					X

The characteristics of the High Plains aquifer in 1980 are summarized by State in table 2. The ranking of the States in 1980 from greatest to least percent of aquifer area are: Nebraska, Texas, Kansas, Colorado, New Mexico, Oklahoma, Wyoming, South Dakota, and Wyoming. The ranking of States in 1980 from greatest to least percentage of aquifer volume is: Nebraska, Texas, Kansas, Colorado, Oklahoma, Wyoming, South Dakota, and New Mexico.

Table 2. Characteristics of the High Plains aquifer in 1980
[Modified from Gutentag and others, 1984]

Characteristic	Unit of measurement	Total	Colorado	Kansas	Nebraska	New Mexico	Oklahoma	South Dakota	Texas	Wyoming
Area underlain by aquifer	Square miles	174,050	14,900	30,500	63,650	9,450	7,350	4,750	35,450	8,000
Percentage of total aquifer area	Percent	100	8.6	17.5	36.6	5.4	4.2	2.7	20.4	4.6
Percentage of each State underlain by aquifer	Percent	—	14	38	83	8	11	7	13	8
Average area-weighted saturated thickness in 1980	Feet	190	79	101	342	51	130	207	110	182
Volume of drinkable water in storage in 1980	Million acre-feet	3,250	120	320	2,130	50	110	60	390	70
Percentage of total volume of drinkable water in storage in 1980	Percent	100	3.7	9.9	65.5	1.5	3.4	1.8	12.0	2.2

HIGH PLAINS WATER-LEVEL MONITORING PROGRAM

Table 3. Number of observation wells measured for the water-level comparison periods—1980-95 and 1994-95

State	Number of observation wells measured	1980 and 1994 and 1995	1995
Colorado	519	571	571
Kansas	965	1,112	1,112
Nebraska	2,000	2,241	2,241
New Mexico	186	211	211
Oklahoma	198	225	225
South Dakota	38	53	53
Texas	1,218	2,218	2,218
Wyoming	21	45	45
High Plains	5,901	7,266	7,266

An extensive network of observation wells is necessary to monitor water levels in the High Plains aquifer (fig. 1). This network consists of many smaller networks of observation wells measured by numerous Federal, State, and local agencies. Local water and natural resource conservation districts are responsible for most of these smaller observation-well networks and the majority of water-level measurements. The total number of wells measured in 1980 and again in 1995 was 5,901; the total number measured in 1994 and again in 1995 was 7,266 (table 3). During the year, the USGS completes the water-level measurements from these local networks and maintains a statewide data base in all of the States in the High Plains region except Texas; in Texas, the Texas Water Development Board maintains the statewide data base. The USGS determines the water level for each measured well that best represents nonpumping conditions and completes the water-level data annually into an aquifer data base.

Water-level measurements are usually made in late winter and early spring when water levels generally represent nonpumping conditions. These measurements normally represent the highest water level during the year. Most observation wells are privately owned irrigation wells. The large diameter and high pumping capacity of these wells make them particularly well suited for monitoring water-level changes because they are less prone to plugging than wells of small diameter and low pumping capacity.

FACTORS AFFECTING WATER-LEVEL CHANGE

Water-level change in the High Plains aquifer results from an imbalance between recharge and discharge. Human activities such as pumping wells and diverting streams have contributed to this imbalance in many parts of the High Plains aquifer, resulting in substantial water-level change through time.

RECHARGE

Precipitation is the principal source of natural ground-water recharge to the High Plains aquifer. Other sources include seepage from streams, canals, and reservoirs and irrigation return flow. Several factors that can affect the proportion of available water that ultimately recharges the aquifer include topography, soil texture and thickness, vegetation or crop type, evapotranspiration, and the lithology and thickness of the unsaturated zone.

Recharge to the High Plains aquifer usually results from the conditions present during the nongrowing season, when evapotranspiration is minimal and soil water can accumulate in the root zone and percolate downward. In the drier western portion of the High Plains region, the recharge process may occur only as an isolated event every several years. In the wetter eastern portion of the High Plains region, the recharge process is more frequent (Dugan and Zelt, in press).

Dugan and Zelt (in press) calculated estimated average annual potential ground-water recharge to the High Plains aquifer from precipitation and irrigation return flow using soil and vegetation characteristics and 1951-80 climatic data. Their recharge estimate ranges from 0.25 to 0.5 inch in the western portion of the High Plains region, where annual precipitation is generally less than 16 inches, to 4-6 inches in the eastern portion of the northern and central High Plains region, where precipitation generally exceeds 24 inches. Expressed as a percentage of average annual precipitation, their recharge rates range from less than 2 percent in the western portion of the High Plains region to more than 17 percent in the eastern portion of the northern High Plains region (McGrath and Dugan, 1993; Dugan and Zelt, in press).

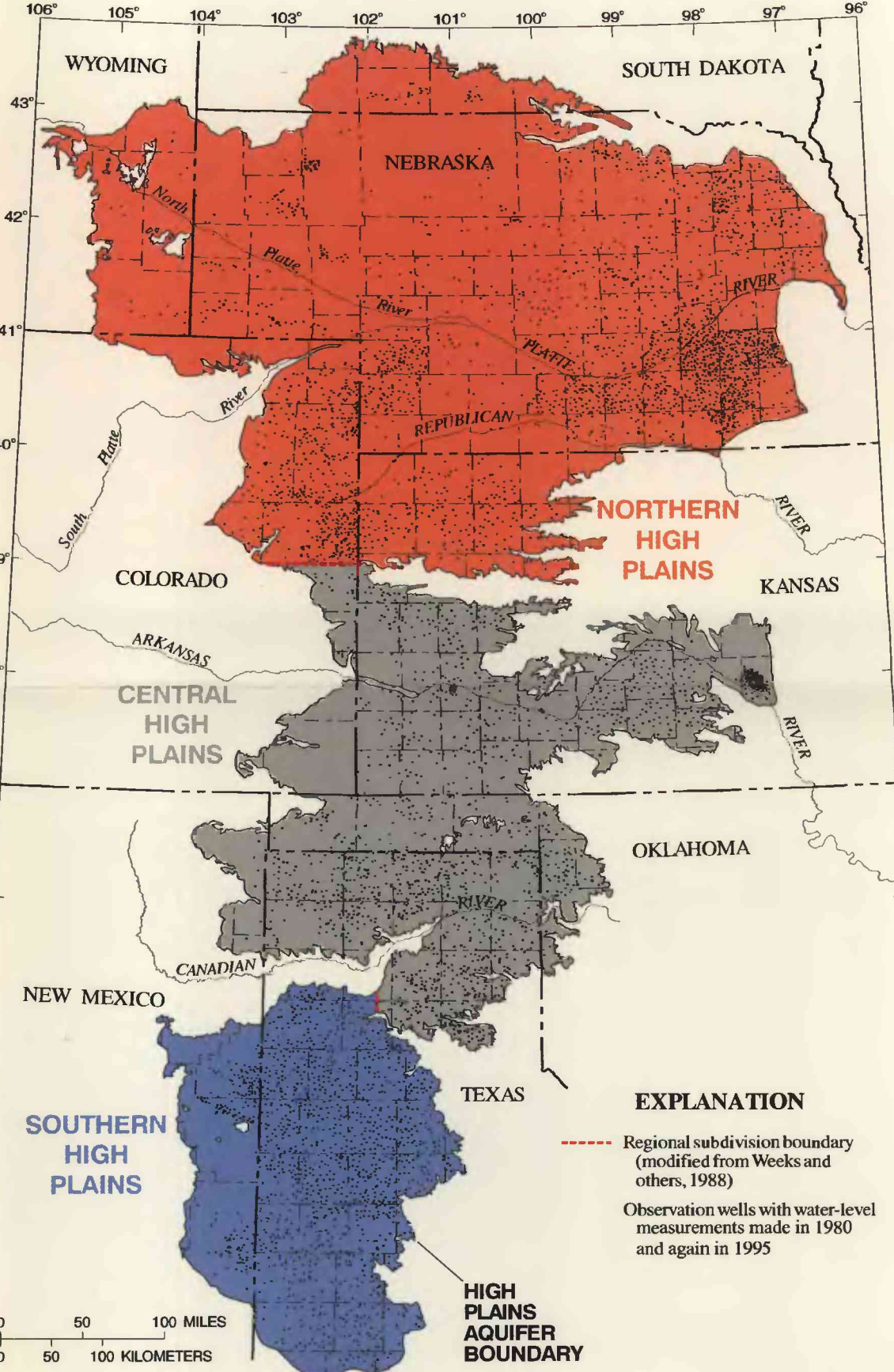


Figure 1. Regional subdivisions of the High Plains aquifer and location of observation wells with water-level measurements in 1980 and in 1995.

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Water-level rises were also observed in portions of the High Plains aquifer in 1980. The largest area of water-level rise was an area about 30 miles long and 6 to 20 miles wide, south of the Platte River in the central portion of the northern High Plains aquifer (fig. 2). In this area, water levels rose more than 10 feet because of seepage losses from surface-water diversions for irrigation and power generation.

WATER-LEVEL CHANGE, 1980 to 1995

Water-level change in the High Plains aquifer from 1980 to 1995 (fig. 3, sheet 2) is based on measurements from 5,901 wells (table 3) and reflects a slightly different pattern of change in comparison to the pattern of change observed from predevelopment to 1980 (fig. 2). Substantial declines have continued in the western portion of the central High Plains aquifer, the northern portion of the southern High Plains aquifer, and the southwestern portion of the northern High Plains aquifer. Some areas that had substantial water-level declines from predevelopment to 1980 in the southern and northern High Plains aquifer, however, had considerably slower rates of decline, or rising water levels, since 1980.

The average area-weighted water level in the High Plains aquifer declined 2.4 feet from 1980 to 1995 compared to 9.9 feet from predevelopment to 1980 (table 5). Considering 1940 as the time of initial irrigation development, the rate of water-level decline decreased from nearly 0.25 foot per year from 1940 to 1980 to 0.16 foot per year from 1980 to 1995. Much of the reduction in the rate of water-level decline since 1980 can be attributed to an average area-weighted water-level rise of 1.8 feet, or 0.12 foot per year, in Nebraska.

The smaller rate of water-level decline for the High Plains aquifer from 1980 can also be attributed, in part, to a much smaller rate of decline in Texas, water levels, which had declined an area-weighted average of 33.7 feet, or 0.84 foot per year, from 1940 to 1980, declined an area-weighted average of 4.8 feet, or 0.32 foot per year, from 1980 to 1995. This decrease in the rate of water-level decline occurred even though total acres irrigated in 1994 from the High Plains aquifer in Texas was 4.05 million acres, which is 0.17 million acres greater than in 1980 (Thein and Heimes, 1987; Texas Water Development Board, 1990). The rate of water-level decline after 1980 slowed in Texas and reversed in some areas in Texas primarily due to decreased pumping as a result of improved irrigation-management practices, (2) a reduction in the number of acres irrigated since the 1970's in those areas in Texas prone to large rates of water-level decline, and (3) annual precipitation averaging as much as 4 inches above normal during 1981-94 (Dugan and Schild, 1992). Some of the areas of water-level rise in the southern High Plains aquifer probably are also associated with the recovery of local cones of depression caused by decreased pumping (Kastner, Schild, and Spahr, 1989).

In contrast, the rate of water-level decline in the High Plains aquifer in Kansas increased after 1980. Considering 1950 as the beginning of irrigation development in Kansas, water levels declined an area-weighted average of 9.9 feet, or 0.33 foot per year, from 1950 to 1980, and an area-weighted average of 7.5 feet, or 0.50 foot per year, from 1980 to 1995. The increase in the rate of decline in Kansas from 1980 to 1995 can be attributed, in part, to the continued increase in irrigation in the late 1970's and early 1980's.

Other factors that appear to have contributed to a reduction in the overall rate of water-level decline in the High Plains aquifer since 1980 include:

1. Average precipitation from 1981 to 1994 was generally greater than normal throughout the High Plains region (fig. 4). Area-weighted average annual precipitation during 1981-94 was 21.82 inches, which is 1.23 inches above, or 106 percent of, normal (table 6).

2. Later phases of irrigation development in the High Plains region shifted geographically from areas of larger potential rates of aquifer depletion to areas of smaller potential rates of depletion. From predevelopment to the 1970's, most irrigated land was in the central and southern High Plains region where recharge tends to be small and CIR large. By 1980, more than half of the 14 million acres irrigated were in the northern High Plains region, mainly in Nebraska, where recharge is generally larger and CIR is smaller than in the central and southern High Plains region (Thein and Heimes, 1987; McGrath and Dugan, 1993; Dugan and Zelt, in press).

3. Significant advances to improve the efficiency of irrigation systems have substantially reduced the ground-water pumping needed to meet CIR. These include the surge-application system with furrow irrigation and low-pressure nozzles on drop tubes to meet center pivot irrigation.

4. Irrigation management practices, including irrigation scheduling based on soil-water conditions and crop growth stages, reuse of irrigation water, and the conversion to crops other than corn or to corn varieties with a smaller CIR, have further reduced ground-water pumping.

5. Water-level declines in some areas prior to 1980 prompted local regulation of ground-water withdrawals for irrigation and development of irrigated land.

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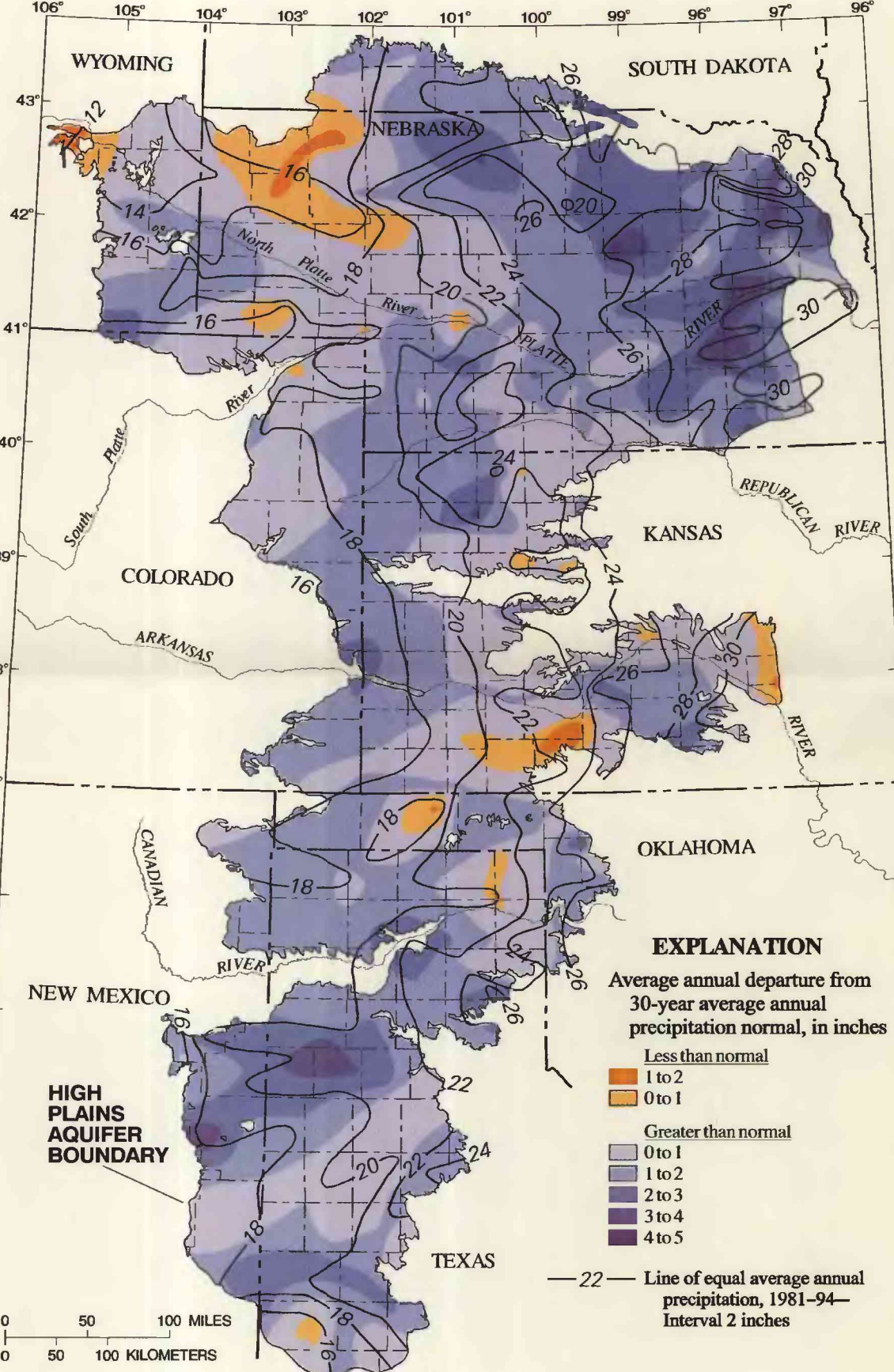


Figure 4. Average annual precipitation, 1981-94, and departure from 30-year normal (1961-90).

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